¹³C/¹²C Partitioning during Synthesis of Na₂Ca(CO₃)₂·2H₂O

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During the synthesis of Na₂Ca(CO₃)₂·2H₂O (pirssonite) from aqueous solutions at 60 and 90 °C, ¹³C is enriched in the solid with respect to the dissolved carbonate species.

Orthorhombic $Na_2Ca(CO_3)_2 \cdot 2H_2O$ 1 may be formed in chemical processes^{1,2} and occurs in nature as the mineral pirssonite.³ Although the determination of the stable isotopic compositions may be useful to resolve the formation conditions of natural and synthetic carbonate solids,⁴ no isotope partition experiments involving 1 have been carried out yet. Therefore, ¹³C/¹²C fractionation during formation of 1 was investigated by synthesis from aqueous solutions.

Compound 1 was synthesized at 60 and 90 °C by the reaction of artificial calcite **2a** [1.2 mmol; $\delta^{13}C = -14.84\%^{\dagger}$] or natural aragonite **2b** [1.2 mmol; δ^{13} C = 5.03‰] with freshly prepared aqueous Na₂CO₃ solutions [39 g; 2.88 mol dm⁻³; δ^{13} C = -5.47%], as in the method of Bury and Redd.¹ The reaction vessels (40 cm³), which were continuously rotated at run temperature in an oven, contained polytetrafluoroethylene-coated metal bars to promote the reaction and to prevent coating of 2 by 1. At the end of each run, the vessel was cooled (H_2O) within a few minutes to room temperature. The solution pH was measured under air-tight conditions, and 1[‡] was separated rapidly from the solution by ultrafiltration under N₂ pressure, washed and dried. The sum of the final dissolved carbonate species 3 was precipitated quantitatively as BaCO₃ 4 in a degassed solution of NaOH and BaCl₂ under N_2

Compound 1 formed in solutions containing dissolved Na⁺ and CO_3^{2-} and 2 according to the overall reaction (1). Since

$$CaCO_3 + 2 Na^+ + CO_3^{2-} + 2 H_2O \rightarrow Na_2Ca(CO_3)_2 \cdot 2H_2O$$
(1)

the sum of the dissolved carbonate species was always in significant excess with respect to 1 and 2, the isotopic compositions of 1 are nearly independent of the initial $CaCO_3$ and the ${}^{13}C/{}^{12}C$ ratios of dissolved carbonate did not change significantly during the interaction process (Table 1). The

 Table 1 Run conditions and final isotopic compositions during synthesis of 1.

Starting CaCO ₃	Conditions	δ ¹³ C (‰) of 1	$\delta^{13}C(\%)^b$ of 4 , measured	δ^{13} C (‰) ^b of 4, calculated
2a	60 °C; 453 h	-2.45	-5.57	-5.69
2a	90 °C; 144 h	-3.35	-5.73	-5.67
2b	90 °C; 360 h	-3.19	a	-5.38

^{*a*} Not determined. ^{*b*} See footnote §.

 Table 2 Measured and calculated fractionation factors involving 1

<i>T/</i> ⁰C	$\alpha(1-3)^a$	$\alpha(1 - CO_3^{2-})$	$\alpha(1 - \text{HCO}_3^-)$	$\alpha(1-\mathrm{CO}_2)$
60	1.0033	1.0033	1.0031	1.0077
90	1.0023	1.0023	1.0023	1.0045

" See footnote §.

measured isotope values of **3** compare well with those obtained from mass balance calculations (Table 1). As **3** is essentially carbonate (CO_3^{2-}) , $\alpha(1-3)$ ¶ virtually equals $\alpha(1 - CO_3^{2-})$ (Table 2). Taking into account the fractionation factors between the dissolved carbonate species and gaseous CO_2 ,⁵ the α -values in Table 2 are obtained. At 60 and 90 °C, ¹³C is enriched in **1** with respect to the dissolved carbonate species considered and gaseous CO_2 , and the degree of fractionation seems to decrease with increasing temperature. The effect of temperature on the isotope fractionation between **1** and the dissolved carbonate species is significantly smaller than for the calculated solid–gas system.

In the temperature range used for the present experiments, isotope fractionation has not hitherto been investigated for hydrated carbonates. However, carbon isotope partitioning has been treated theoretically in several studies⁶ for the system calcite–CO₂. In the most recent work, Chacko *et al.*⁶ obtained α (calcite-CO₂) values of 1.0071 and 1.0047 at 60 and 90 °C, respectively, which are similar to the present experimental results for **1** (Table 2).

The IR spectra were obtained by Dr P.-L. Gehlken.

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Footnotes

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‡ 1 was identified by powder XRD and IR spectroscopy.

§ The isotopic composition of 3 equals that of 4.

 $\P \alpha(\mathbf{A}-\mathbf{B}) = [\delta^{13}\mathbf{C}_{\mathbf{A}} + 1000]/[\delta^{13}\mathbf{C}_{\mathbf{B}} + 1000].$

 $\|$ A similar behaviour seems to occur for ¹⁸O/¹⁶O of the carbonate ion group of 1.

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