

Static SIMS Studies of Reactions on Mimics of Polar Stratospheric Clouds II: Low-Temperature, Low-Pressure Interactions of Cl₂ and Cl₂O with Solid Ice Films

Heather A. Donsig, Dawn Herridge, and John C. Vickerman*

Surface Analysis Research Centre, Department of Chemistry, UMIST, P.O. Box 88, Manchester M60 1QD, U.K.

Received: September 15, 1997; In Final Form: December 10, 1997

The interactions of molecular chlorine, dichlorine monoxide, and hypochlorous acid with solid ice films have been examined. At 90–100 K, chlorine can be made to condense as a molecular solid, forming a discrete layer on top of an existing film of solid H₂O. Interaction between the Cl₂ and the water molecules is minimal. At temperatures above approximately 130 K, the Cl₂ and water react to produce a mixed film of water, hypochlorous acid, and solvated HCl. As with chlorine, the dichlorine monoxide did not react with water at 90 K. However, there was evidence of a hydrogen-bond-like interaction between the Cl₂O and H₂O molecules. Upon warming to 120 K or higher, the two reacted to produce a mixed film of hypochlorous acid and water. The reactions that occurred at the higher temperatures are of interest because of the well-documented influence of polar stratospheric clouds, which consist of solid water or water/nitric acid particles, on the release of active chlorine species into the stratosphere.

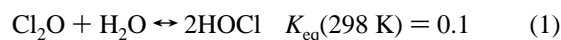
Introduction

The discovery that the clouds that form every winter over the Antarctic play an important role in ozone depletion has prompted numerous studies of the possible chemistry involved. These clouds are made up of tiny particles of condensed water/nitric acid mixtures or of water alone condensed around a water/nitric acid core. In the presence of these tiny cloud particles, the conversion of relatively inert chlorine-containing species into molecular chlorine is greatly enhanced.

The production and release of molecular chlorine is a crucial step in the ozone-depletion cycle, since the Cl₂ photolyzes readily in sunlight to provide a steady source of chlorine radicals. Thus, the heterogeneous chemistry of gaseous chlorine-containing species interacting with a solid icy surface is of great interest to atmospheric chemists. In this study, we use static secondary-ion mass spectrometry (SSIMS) to study the interactions of molecular chlorine, of dichlorine monoxide, and of hypochlorous acid with films of condensed water grown under ultrahigh vacuum (UHV) conditions. Our purpose was to ascertain the nature of the interactions of these species with ice, along with the nature of any reaction products or reactive intermediates formed.

Previous studies have used IR spectroscopy and temperature-programmed mass spectrometry (TPMS) to study the interactions of chlorine-containing species with ice films. Reflection-absorption IR spectroscopy (RAIRS) in particular is useful for studying reactions at surfaces. For example, RAIRS has been used to follow the reaction between molecular chlorine and water co-deposited onto a cooled, inert substrate.¹ At 180 K, under UHV conditions, neither water nor molecular chlorine condensed individually. When both were introduced simultaneously, the IR spectrum observed was thus attributed to the condensation of a reaction product. The IR spectrum of the condensed product contained a strong, sharp absorption band, blue-shifted from that due to condensed water. This band is apparently characteristic of water of hydration, leading to speculation that it arose from a [H₂OCl]⁺ ion.

The same study examined the reactions of dichlorine monoxide (Cl₂O) on ice. Although it is not a major constituent of atmospheric chlorine, dichlorine monoxide in the gas phase reacts with water in an equilibrium reaction to produce hypochlorous acid:²



This reaction, and particularly the hypochlorous acid that is formed, is of interest due to the host of reactions that could take place on the watery surfaces of polar stratospheric cloud particles.

The RAIRS spectrum of Cl₂O and H₂O codeposited at 80 K from the gas phase clearly showed stretching modes from pure molecular Cl₂O, along with weak bands attributed to solid–solid interactions between the water and the dichlorine monoxide. A spectrum of solid Cl₂O condensed on top of a solid ice film was similar but lacked the weak solid–solid interaction bands. Premixing Cl₂O and water in the gas phase, then condensing the mixture onto the cooled substrate gave a mixed H₂O–Cl₂O–HOCl film, with the HOCl showing a large degree of hydrogen bonding. Heating to 140 K removed all traces of residual Cl₂O.

The IR experiments proved very informative and laid an important foundation in the study of these systems. However, no one technique is ideal for providing definitive, unambiguous answers to all the questions one may wish to address. Static secondary-ion mass spectrometry (SSIMS) was adopted as a highly sensitive technique that is capable of providing detailed chemical information about species as they exist in situ, on the surface of the film.³ In static SIMS, secondary particles (anions, cations, neutrals, and electrons) are ejected from the surface as a result of bombardment by a high-energy primary ion beam. When the primary ion dose is kept low (less than 10¹³ ions/cm²), damage to the sample during analysis can be minimized. The sputtered secondary ions are then detected by a mass spectrometer. By use of time-of-flight analysis, very high

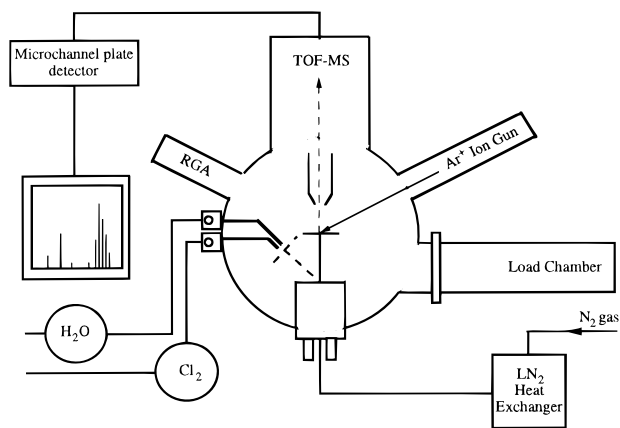


Figure 1. Schematic of the secondary-ion mass spectrometer and sample-dosing apparatus: TOF-MS, time-of-flight mass spectrometer, detector for the secondary ions; RGA, quadrupole mass spectrometer used for residual-gas analysis.

transmissions can be obtained for these secondary ions, making SIMS an extremely sensitive technique for detecting trace quantities of ionizable species from the surface.

Little data are available on the characteristic SIMS spectra of condensed systems such as these. Thus, it was necessary to establish a base of standard "reference" spectra for simple systems to aid in the interpretation of the more complex systems. This was done for pure water, for molecular chlorine, and for dichlorine monoxide by condensing each reagent individually onto an inert gold substrate at 90–100 K. Mixed systems of relevance to atmospheric chemistry, such as Cl₂/water or Cl₂O/water, were then studied for comparison.

In the stratosphere, cloud particles exist at temperatures of 185 to 210 K. We are therefore more interested in how these systems react at higher temperatures. Under UHV conditions, condensed water begins to desorb rapidly by 150 K,⁴ and the reactants tended to desorb at even lower temperatures.⁵ Thus, for the work at higher temperatures, it was sometimes deemed appropriate to establish a constant backpressure of water, the reactant, or both (to a total pressure of 10⁻⁵ mbar or less) in order to maintain a film of reasonable thickness. Doing so produced conditions that were more representative of the stratosphere, where the total pressures are in the range 1–100 mbar (depending on the exact altitude). This means that the surfaces of stratospheric cloud particles exist in a state of dynamic flux, with material constantly desorbing and recondensing.⁴ By raising the temperature until material began to desorb rapidly, then establishing a sustaining backpressure of the relevant species, we were able to simulate these conditions in our vacuum chamber.

Experimental Section

The experiments were performed using a Kratos time-of-flight secondary-ion mass spectrometer (TOF-SIMS) equipped with a special gas-dosing and -handling system, a temperature-controlled sample manipulator, and a VG micromass quadrupole mass spectrometer for residual gas analysis (Figure 1). The TOF-SIMS uses 10 kV Ar⁺ ions as the primary source, with typical ion doses during data acquisition being on the order of 2 × 10¹² ions/cm². Details of the system have been described elsewhere.⁶

The chlorine (99.5% purity) was used as purchased from Aldrich. The Cl₂O was synthesized by passing a stream of Cl₂ in nitrogen (zero grade, 99.995%) through a glass column packed with dry, yellow mercuric oxide (Aldrich, 99+%)

dispersed on a bed of glass beads.⁷ The product was trapped under liquid nitrogen, then purified by pumping briefly under dry ice/acetone to remove residual Cl₂.

Water, Cl₂, and Cl₂O were introduced into the vacuum chamber from separate reservoirs in the glass dosing line using a set of leak valves. Hypochlorous acid (HOCl) was synthesized at room temperature in one of the reservoirs by mixing gaseous Cl₂O and H₂O. The resulting Cl₂-H₂O-HOCl mixture was then introduced through a single leak valve. Inside the vacuum chamber, a set of 3-mm diameter Teflon tubes were used to direct the reagents. The sample stub could be rotated out of the analysis position to one directly in front of the nozzles, causing the relatively large amounts of the material to be selectively deposited onto the gold substrate (dosing). Alternatively, when smaller amounts were required, or to help maintain a film during analyses at higher temperatures (140 K or higher), material could be backdosed by opening the leak valves while leaving the sample stub in the analysis position.

Under the dosing conditions normally employed, we expect to form an ice film consisting of several hundred to a thousand layers. Except where otherwise noted, the water films were then thermally annealed to 150 K, then cooled back to the temperature of the experiment before introducing the second component. This should cause the ice film, which is almost certainly amorphous as dosed, to crystallize.

Results and Discussion

Molecular Chlorine at T = 90 K. Spectra were obtained at 90 K for both neat Cl₂ condensed onto a gold substrate and for Cl₂ condensed on ice. The spectra of pure Cl₂ on gold were very similar to the Cl₂/ice spectra, which are shown in Figure 2. In both cases the positive-ion SIMS spectra consisted of chlorine-cluster ion peaks (Cl_m⁺, 1 ≤ m ≤ 10 or higher) plus (H₂O)_nH⁺ and (H₂O)_n⁺ cluster peaks. The last two families are typical of systems in which condensed water is present, and in the spectrum of "neat" Cl₂ on gold, they indicate the presence of trace quantities of residual water.

The negative-ion SIMS spectrum of condensed Cl₂ on gold similarly consisted of a series of Cl_m⁻-cluster ion peaks (1 ≤ m ≤ 10 or higher). The O⁻ and OH⁻ peaks arise from condensed water and are very much smaller in the spectrum of neat Cl₂ on gold than in the Cl₂/ice spectrum. The negative-ion Cl_m⁻ clusters give much more intense signals than the corresponding positive-ion clusters, which is reasonable, since chlorine is highly electronegative. In both the positive- and negative-ion spectra, the m = 1 peak was the most intense, with the intensities decreasing uniformly as m increased.

It is important to make the comparison between the Cl₂/ice and neat Cl₂ spectra. The Cl₂/ice spectrum was a simple superposition of the corresponding neat Cl₂ and water spectra. No new peaks were observed. The (H₂O)Cl⁻ ion was present in both cases in the "neat" spectrum, presumably forming from the trace quantities of residual water that condensed along with the Cl₂. In general, there were no significant, systematic shifts in the relative peak intensities. Changes of this sort would be expected if a reaction were occurring, allowing us to conclude that Cl₂ does not chemically interact with ice at 90–100 K.

Dichlorine Monoxide at T = 90 K. The experiment was repeated for Cl₂O, again at temperatures of 90–100 K. The Cl₂O spectra were more complex than the Cl₂ spectra, although the major peaks were similar. The characteristic patterns for mixtures of the ³⁵Cl and ³⁷Cl isotopes simplified interpretation of the spectra. Figure 3 shows typical spectra of neat Cl₂O condensed on gold, which in general gave clearer signals than

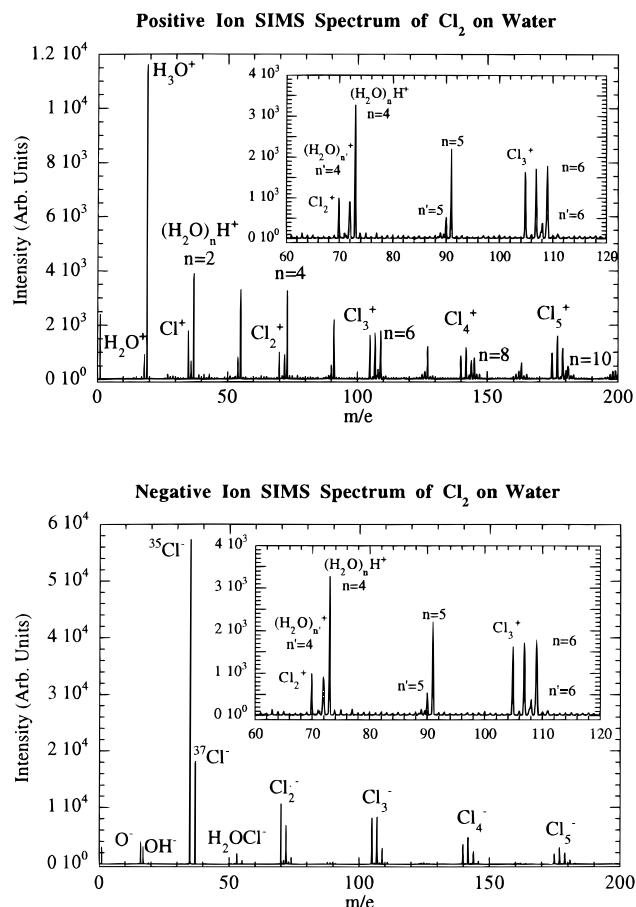


Figure 2. Positive- and negative-ion SIMS spectra of chlorine condensed on ice. Water was dosed to 2×10^{-7} mbar for 30 s at 95 K, and the sample was unannealed. Chlorine was dosed to 4×10^{-8} mbar for 30 s at 95 K. Spectra recorded at 95 K.

did the $\text{Cl}_2\text{O}/\text{ice}$ films. Trace quantities of water were codeposited along with the Cl_2O . The water peaks are clearly evident in the positive-ion spectrum, where the most prominent peaks are the $(\text{H}_2\text{O})_n\text{H}^+$ -cluster ions. The unprotonated $(\text{H}_2\text{O})_n^+$ clusters are less prominent. As with the Cl_2 , the spectra of Cl_2O on ice were similar to the spectra of neat Cl_2O (minimal water present).

Numerous minor peaks appeared in the positive- and negative-ion spectra. These are itemized in Tables 1 and 2, which together give a comprehensive list of peaks that appeared in the various systems studied. To facilitate comparisons between spectra, the intensities in the negative-ion spectra are tabulated as a percentage of the total over all peaks. In the positive-ion table, the water-cluster peaks for $n > 2$ have been omitted from the total to avoid obscuring the important chlorine-containing species.

For Cl_2O (Figure 3) and Cl_2O on ice at 90 K (Tables 1 and 2), it is worth noting the presence of the molecular ion peak in both the positive- and negative-ion spectra. In fact both Cl_2O^+ ($m/e = 86$) and the protonated molecular ion Cl_2OH^+ ($m/e = 87$) appeared in the positive-ion spectrum, with the analogous Cl_2O^- and Cl_2OH^- ions appearing in the negative-ion spectrum. In both cases they appeared as relatively minor peaks and were most readily observed when the Cl_2O overlayer was very pure. However, that the parent ions appeared in both the positive- and negative-ion spectra suggests that the Cl_2O molecules remain intact and condense on top of, or cocondense with, the water molecules at 90–100 K.

Although new peaks that might indicate the formation of

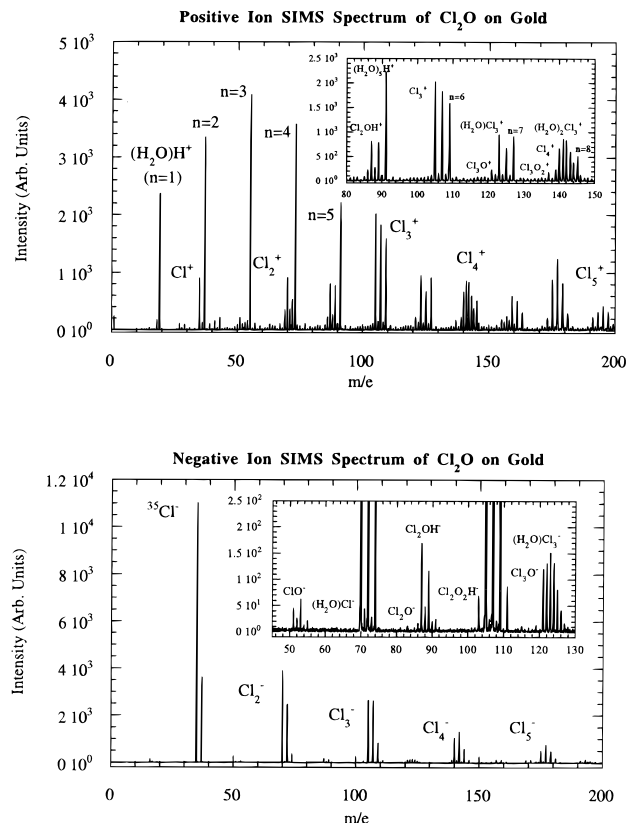


Figure 3. Positive- and negative-ion SIMS spectra of dichlorine monoxide condensed onto a clean gold substrate. Cl_2O was backdosed to 5×10^{-8} mbar at 100 K. Spectra were recorded at 100 K.

reaction products did not appear in either case, the positive-ion spectra show an interesting and unique feature that bears thinking about. We have observed in general that low-temperature positive-ion spectra, even those of the “neat” compounds condensed on gold, contain two sets of water-cluster peaks, namely, the $(\text{H}_2\text{O})_n^+$ and $(\text{H}_2\text{O})_n\text{H}^+$ families. In the SIMS spectrum of pure condensed water, the dominant peak in each is the $n = 1$ peak. The cluster intensities then decrease as a function of n , with the exception of $n = 4$ where a local maximum may appear. This enhanced intensity for water clusters with $n = 4$ has been observed previously^{8–10} and can be attributed to the exceptional stability of the $(\text{H}_2\text{O})_4\text{H}^+$ ion. The existence of this and other exceptionally stable clusters (with “magic numbers” at $n = 4, 21,$ and 28) is widely acknowledged, and the phenomenon has been well studied.^{11–14}

In our experiments, the introduction of HCl ,⁶ molecular chlorine, or other adsorbates in general¹⁵ did not perturb the water-cluster distribution. The local maximum at $n = 4$ was not uniformly evident. Its appearance may depend on the exact surface density or distribution of the water molecules, since it was most pronounced in the pure water spectra. However, when Cl_2O was present, the distribution was markedly different. The maximum overall intensity was shifted to the $n = 3$ cluster, rather than occurring at $n = 1$ as observed in all other instances. The difference is illustrated in Figure 4, which plots the relative peak intensities of the $(\text{H}_2\text{O})_n\text{H}^+$ ions as a function of cluster number (n) at 90–100 K for pure water, for Cl_2O alone and on ice, and for related systems studied in our lab. In each case the most intense water-cluster peak intensity was normalized to unity. The maximum at $n = 3$ in the Cl_2O spectra is clearly visible and is in marked contrast to the distributions observed in the other systems.

This difference is clearly significant and may arise from the

TABLE 1: Integrated Peak Intensities from Positive-Ion SIMS Spectra^a

<i>m/e</i> , ion	HCl/ice 90 K	Cl ₂ /ice 90 K	Cl ₂ O/ice 100 K	Cl ₂ O/H ₂ O 140 K	HOCl/H ₂ O 130 K	Cl ₂ /H ₂ O 130 K
18, H ₂ O ⁺	3.8	3.8	0.7	5.2	3.4	4.0
19, H ₃ O ⁺	79.8	45.9	15.7	72.2	81.5	79.2
35, Cl ⁺		7.3	6.6		0.1	
36, (H ₂ O) ₂ ⁺	0.6	2.9	0.4	1.3	0.5	0.5
37, (H ₂ O) ₂ H ⁺	15.2	17.3	26.1	18.0	13.8	16.3
51, ClO ⁺	0.2	0.2	2.8	0.6	0.1	
52, HOCl ⁺		0.2	0.7		trace	
53, H ₂ OCl ⁺	0.2	0.4	1.2	1.2	0.2	
67, ClO ₂ ⁺			2.8		trace	
69, H ₂ OClO ⁺	0.1		2.6	0.8	trace	
70, Cl ₂ ⁺		4.5	5.3			
71, HCl ₂ ⁺		0.4	2.6	0.1	0.4	
85, H ₂ OCl ₂ O ⁺			0.5			
86, Cl ₂ O ⁺			2.2			
87, HCl ₂ O ⁺			3.5			
88, H ₂ OCl ₂ ⁺			2.6			
89, H ₂ OHCl ₂ ⁺	0.1	0.6	4.0	0.4		
102, Cl ₂ O ₂ ⁺			1.6			
105, Cl ₃ ⁺		7.2	10.9			
140, Cl ₄ ⁺		4.2	3.2			
175, Cl ₅ ⁺		5.1	4.0			
19/18, H ₃ O ⁺ /H ₂ O ⁺	21.0	12.1	22.4	13.9	24.0	19.8

^a All intensities are normalized, expressed as a percentage of the total counts from all peaks in a given spectrum, to facilitate comparisons among the various systems: for HCl/ice, 90 K,⁶ water dosed to 6.6×10^{-7} mbar for 50 s, annealed 10 min at 150 K, HCl backdosed to 6.2×10^{-8} mbar for 35 s; for Cl₂/ice, 90 K, as per Figure 2; for Cl₂O/ice, 100 K, water dosed to 1×10^{-7} mbar for 120 s, annealed 10 min at 150 K, Cl₂O dosed to 1.5×10^{-7} mbar for 135 s; for Cl₂O/H₂O, 140 K, water dosed to 2.3×10^{-7} mbar for 120 s, annealed 10 min to 150 K, Cl₂O backdosed to 1×10^{-7} mbar (the Cl₂O reacts, producing a film of hypochlorous acid (see text)); for HOCl/H₂O, 130 K, gaseous Cl₂O and H₂O premixed at room temperature, dosed to 1×10^{-7} mbar (total) for 60 s at 90 K, heated slowly to 130 K; for Cl₂/H₂O, 140 K, water dosed to 3.5×10^{-7} mbar for 60 s at 90 K, annealed 10 min to 150 K, Cl₂ dosed to 3.0×10^{-7} mbar for 60 s at 80 K, heated slowly to 130 K.

TABLE 2: Integrated Peak Intensities from Negative-Ion SIMS Spectra^a

<i>m/e</i> , ion	HCl/ice 90 K	Cl ₂ /ice 90 K	Cl ₂ O/ice 100 K	Cl ₂ O/H ₂ O 140 K	HOCl/H ₂ O 130 K	Cl ₂ /H ₂ O 130 K
16, O ⁻	20.6	15.1	4.2	13.5	15.2	18.3
17, OH ⁻	8.6	9.4	0.4	13.1	9.4	12.4
35, Cl ⁻	69.6	71.1	78.3	65.2	71.6	67.0
51, ClO ⁻	0.1	0.1	0.2	0.3	0.1	0.1
52, ClOH ⁻	0.1	0.1		0.3	0.1	0.1
53, H ₂ OCl ⁻	0.6	1.9	0.1	5.3	1.9	1.5
70, Cl ₂ ⁻	0.2	1.0	6.1	0.5	1.0	0.2
71, HCl ₂ ⁻	0.2	0.7	0.1	1.1	0.3	0.3
86, Cl ₂ O ⁻			0.2			
87, HCl ₂ O ⁻			0.4			
88, H ₂ OCl ₂ ⁻	trace	0.1	0.1	0.2	trace	
89, H ₂ OHCl ₂ ⁻	trace	0.1	0.2	0.5	trace	0.1
105, Cl ₃ ⁻		0.3	4.3		0.3	
121, Cl ₃ O ⁻			1.0			
122, Cl ₃ OH ⁻			0.2			
137, Cl ₃ O ₂ ⁻			0.6			
139, H ₂ OCl ₃ O ⁻			0.8			
140, Cl ₄ ⁻		0.1	1.5		0.1	
175, Cl ₅ ⁻		trace	1.1			
191, Cl ₅ O ⁻			0.3			
71/70, HCl ₂ ⁻ Cl ₂ ⁻	1.0	0.70	0.02	2.1	0.30	1.5

^a All intensities are expressed as a percentage of the total counts from all peaks in a given spectrum: for HCl/ice, 90 K, water dosed to 5×10^{-7} mbar for 30 s, unannealed, HCl backdosed continuously to 1.5×10^{-7} mbar; for Cl₂/ice, 90 K, as per Figure 2; for Cl₂O/ice, 100 K, as per Table 1; for Cl₂O/H₂O, 140 K, as per Table 1; for HOCl/H₂O, 130 K, as per Table 1; for Cl₂/H₂O, 140 K, water dosed to 4.0×10^{-7} mbar for 60 s at 83 K, annealed 10 min at 150 K, Cl₂ dosed to 3.0×10^{-7} mbar for 60 s at 90 K, system heated slowly to 140 K, Cl₂ continuously backdosed to 9.3×10^{-8} mbar at 140 K.

way dichlorine monoxide and water molecules interact at this temperature. In a film of pure water ice, the molecules interact by hydrogen bonding. Each δ^+ hydrogen in one water molecule is capable of interacting with the δ^- oxygen of an adjacent molecule to form an extensive, long-range hydrogen-bonded network. One would expect the extent of hydrogen bonding to influence the SIMS spectrum, for instance, by affecting the observed cluster distribution. Systematic changes have in fact

been observed in the distributions from water and from C₁ to C₈ alcohols.⁹

Consider the effect of introducing Cl₂O to a film of pure water ice. The Cl₂O molecule can be thought of as a substituted analogue of water, since the relative electronegativities are such that the oxygen atom is polarized δ^- while the chlorines are δ^+ . This means that the Cl₂O molecules can interact as the water molecules themselves do (Figure 5), inserting themselves

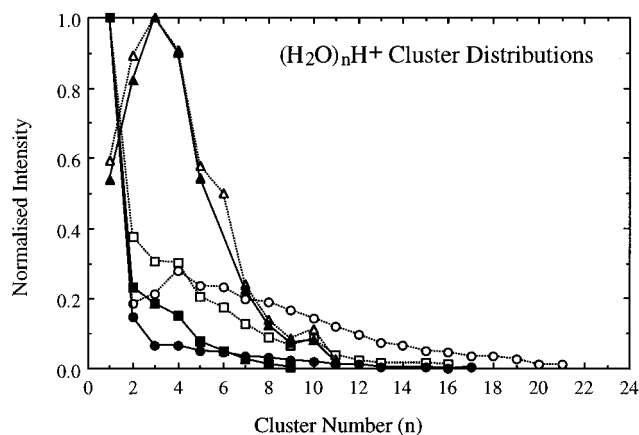


Figure 4. Comparison of $(\text{H}_2\text{O})_n\text{H}^+$ -cluster intensities as a function of cluster number n for various systems at $T \approx 90$ K: (○) pure ice; (●) HCl on ice;⁶ (□) Cl_2 on ice; (△) Cl_2O on gold; (▲) Cl_2O on ice; (■) ClONO_2 on ice.¹⁵ In each case, the maximum $(\text{H}_2\text{O})_n\text{H}^+$ intensity has been normalized to unity.

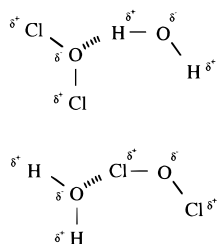


Figure 5. Schematic showing the hydrogen bonding or hydrogen-like bonding that is possible between water molecules and dichlorine monoxide.

and interrupting the long-range network of tetrahedrally coordinated pure water. This could explain the observed perturbation of the water-cluster ion distribution in low-temperature systems containing Cl_2O , as well as the presence of the H_3O^+ analogues H_2ClO^+ and HCl_2O^+ . The other species studied, such as HCl, Cl_2 , ClONO_2 , etc., do not have the same effect presumably because they do not “mimic” water molecules in the same way. Among the molecules studied in our lab,^{6,15} Cl_2O is unique in interacting this way with ice at 90–100 K.

Cl_2O and Water at Higher Temperatures. Temperatures in the stratosphere are typically much greater than 90–100 K, and it is at these higher temperatures that we might expect reactions to occur, either between coadsorbed species or between an adsorbate and the ice film itself. Naturally, the spectra obtained from mixed systems, or in systems where a reaction has occurred, become more complex. These spectra are best interpreted by starting at a low temperature, then monitoring the changes observed as the temperature is increased.

By comparison of the Cl_2O /water spectra at 90 and at 140 K (Tables 1 and 2), it is immediately clear that raising the temperature caused the features associated with the Cl_2O to disappear. Most obvious were the loss of the Cl_m^- peaks for $m > 2$ in the negative-ion spectrum, the loss of all Cl_m^+ peaks in the positive-ion spectrum, and the shift of the $(\text{H}_2\text{O})_n\text{H}^+$ -cluster ion distribution back to “normal” (maximum intensity for $n = 1$). As with the other mixed systems, there was no local maximum for $n = 4$. The relative intensities of the water peaks increased in both the positive- and negative-ion spectra, indicating an increased proportion of water at the surface.

The changes of the sort described above would result from the desorption of Cl_2O , which occurs at 120 K.¹ Certain other changes occurred at this stage, which seem to indicate that the

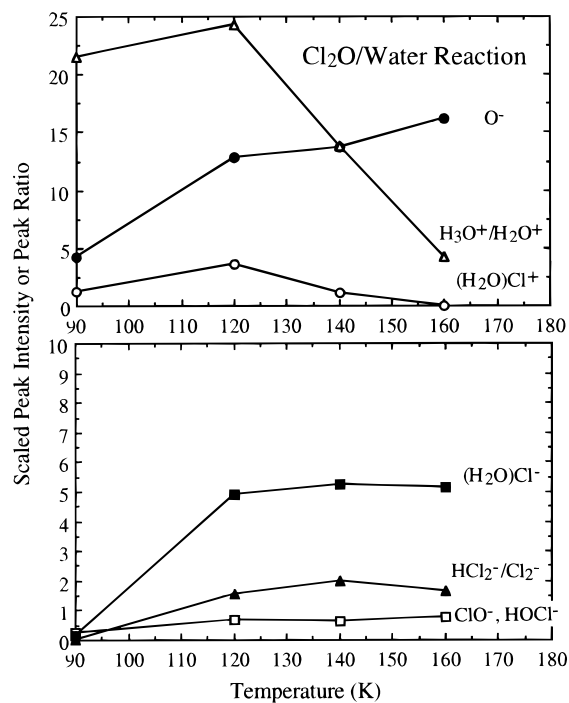


Figure 6. Highlights of the changes that occur in the Cl_2O /ice system as the temperature is raised from 90 to 160 K: (△) $\text{H}_3\text{O}^+/\text{H}_2\text{O}^+$ ratio; (●) O^- normalized intensity; (○) $(\text{H}_2\text{O})\text{Cl}^+$ normalized intensity; (■) $(\text{H}_2\text{O})\text{Cl}^-$ normalized intensity; (▲) $\text{HCl}_2^-/\text{Cl}_2^-$ ratio; (□) ClO^- , HOCl^- normalized intensities. All normalized intensities were scaled as a percentage of the total counts from all peaks in a given spectrum. Cl_2O was dosed to approximately 1×10^{-6} mbar for 300 s at 90 K.

Cl_2O does not merely desorb from the film. Figure 6 highlights the important changes. For instance, as the temperature was raised, there was a decrease in the $\text{H}_3\text{O}^+/\text{H}_2\text{O}^+$ ratio in the positive-ion spectrum, an increase in the relative importance of the O^- and $(\text{H}_2\text{O})\text{Cl}^-$ ions in the negative-ion spectrum, and a dramatic increase in the $\text{HCl}_2^-/\text{Cl}_2^-$ ratio. At temperatures of 130 K and higher, Cl_2O was backdosed in order to maintain the film. By 160 K, it was necessary to backdose the water as well. The changes seen in the film suggest the simultaneous occurrence of a reaction between Cl_2O and water, with products that remain condensed at temperatures of 140 K or higher.

Water and Cl_2O are known to react in the gas phase to produce an equilibrium mixture of Cl_2O , HOCl , and water. It is therefore interesting to compare the high-temperature spectra of Cl_2O adsorbed on a pregrown ice film with spectra of Cl_2O that has been premixed with water, then condensed onto the gold substrate. Table 3 shows data from Cl_2O and water that were premixed in roughly equal proportions at room temperature, then condensed at 90–100 K. The initial spectra taken at this temperature show some of the features associated with low-temperature Cl_2O on ice. These include the presence of Cl_m^- clusters for $m > 2$ in both the positive- and negative-ion spectra, the presence of the $(\text{H}_2\text{O})\text{Cl}^+$ ion in the positive-ion spectrum, and the low $\text{HCl}_2^-/\text{Cl}_2^-$ ratio (0.082) in the negative-ion spectrum.

Warming the film to 130 K caused the expected loss of the Cl_2O features and an increase in the $\text{H}_3\text{O}^+/\text{H}_2\text{O}^+$ ratio, along with the increase in the $\text{HCl}_2^-/\text{Cl}_2^-$ ratio, and O^- and $(\text{H}_2\text{O})\text{Cl}^-$ intensities observed upon warming straight Cl_2O on ice. Again, these changes are consistent with the loss of Cl_2O and the formation of a reaction product or products, which remain condensed. At this temperature, introducing a backpressure of the $\text{Cl}_2\text{O}/\text{H}_2\text{O}/\text{HOCl}$ equilibrium mixture should result in no

TABLE 3: Highlight of Changes in the Spectra from a Cl₂O/H₂O/HOCl Mixture Formed at Room Temperature by Premixing Gaseous Cl₂O and H₂O in the Gas-Dosing/Preparation Line^a

	Positive Ions				
	<i>m/e</i> = 18 H ₂ O ⁺	<i>m/e</i> = 19 H ₃ O ⁺	<i>m/e</i> = 53 (H ₂ O)Cl ⁺	<i>m/e</i> = Cl _{<i>m</i>} ⁺ <i>m</i> ≤ 4	19/18 H ₃ O ⁺ /H ₂ O ⁺
100 K	4.5	67.4	0.5	5.6	14.9
130 K	3.4	81.5	0.2	0.3	24.0
backdosing, 130 K	3.2	81.5	0.1	25.5	
	Negative Ions				
	<i>m/e</i> = 16 O ⁻	<i>m/e</i> = 35 Cl ⁻	<i>m/e</i> = 53 (H ₂ O)Cl ⁻	Cl _{<i>m</i>} ⁻ <i>m</i> ≤ 4	71/70 HCl ₂ ⁻ /Cl ₂ ⁻
100 K	5.0	73.2	1.6	88.1	0.08
130 K	15.2	71.6	1.9	73.0	0.30
backdosing, 130 K	22.0	60.9	2.3	61.2	2.2

^a All intensities are expressed as a percentage of the total counts from all peaks in a given spectrum. Mixture was dosed as per HOCl/H₂O in Table 1, slowly heated to 130 K. Backdosing was to a total pressure of 5×10^{-7} mbar.

further condensation of Cl₂O. Opening the leak valve did produce further changes, which must therefore be attributed to the condensation of mixed hypochlorous acid (HOCl) and water. The changes appeared primarily in the negative-ion spectrum, where the O⁻ and (H₂O)Cl⁻ increased in both absolute and relative intensity, while the HCl₂⁻/Cl₂⁻ ratio jumped from 0.28 to 2.15. These are the traits that appeared when pure Cl₂O condensed on ice was warmed to 140 K. The definite similarities between the two systems suggest that upon warming, the condensed Cl₂O and water react to produce hypochlorous acid (HOCl) at the ice surface.

Cl₂ and Water at Higher Temperatures. The conversion of Cl₂O to hypochlorous acid at the surface upon warming the film is not especially surprising. Of greater interest is a possible reaction between condensed molecular chlorine and ice as the temperature is raised. Tables 1 and 2 highlight the differences between Cl₂/water spectra at 90 and at 130 K. Loss of the Cl_{*m*}[±] peaks, coupled with a rise in the background partial pressure of Cl₂ (monitored on the quadrupole mass spectrometer used for residual-gas analysis), indicated desorption of at least some Cl₂. However, as in the case of Cl₂O, other changes appeared that seem to indicate that a reaction occurs as well. These changes are illustrated by a series of positive- and negative-ion spectra, recorded at temperatures between 100 and 165 K. Certain key quantities were plotted as a function of temperature, with the Cl₂ backdosed as necessary to maintain a constant backpressure of approximately 2×10^{-7} mbar (Figure 7). As indicated by the graphs, the H₃O⁺/H₂O⁺ ratio increased with temperature, with the biggest change coming between 110 and 130 K. The O⁻ relative intensity showed an abrupt increase at temperatures above 130 K. The relative intensities of the ClO⁻, HOCl⁻, and (H₂O)Cl⁻ ions all increased steadily with temperature, as did the HCl₂⁻/Cl₂⁻ ratio.

Certain similarities exist between these graphs and the graphs from the Cl₂O/water system (Figure 6), suggesting that similar reactions involving HOCl may be occurring in both systems. However, there are significant differences, including a decrease in the H₃O⁺/H₂O⁺ ratio from Cl₂O/water above 120 K (Figure 6), a decrease that was not observed when Cl₂O was backdosed at 130 K (Table 3). For Cl₂/water, with a constant Cl₂ backpressure of approximately 2×10^{-7} mbar, the ratio remains high. A high H₃O⁺/H₂O⁺ ratio has previously been associated with a high concentration of acidic protons within the film.⁶ There is also a slight falloff of the HCl₂⁻/Cl₂⁻ ratio for the Cl₂O/water at 160 K, as opposed to a steady rise in the case of Cl₂/water. These changes may signify the presence of HOCl plus water in the Cl₂O/water case, and both HOCl and HCl plus

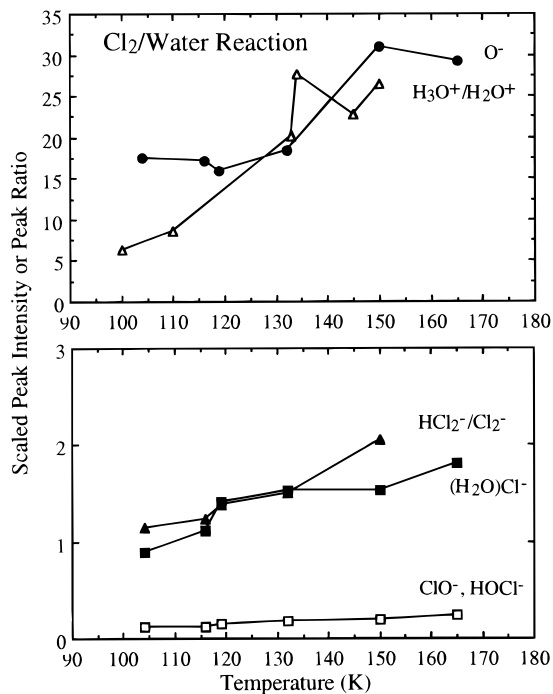
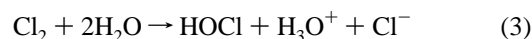
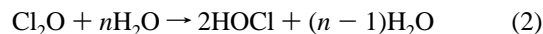


Figure 7. Highlights of the changes that occur in the Cl₂/ice system as the temperature is raised from 105 to 165 K: (Δ) H₃O⁺/H₂O⁺ ratio; (●) O⁻ normalized intensity; (■) (H₂O)Cl⁻ normalized intensity; (▲) HCl₂⁻/Cl₂⁻ ratio; (□) ClO⁻, HOCl⁻ normalized intensities. All normalized intensities were scaled as a percentage of the total counts from all peaks in a given spectrum. Cl₂ background partial pressure was held approximately constant at 2×10^{-7} mbar.

water in the case of Cl₂/water, with net reactions that proceed as follows:



Reaction 2 is consistent with the known gas- or liquid-phase chemistry of Cl₂O and water. It is also consistent with the hydrogen bondlike interactions between Cl₂O and water (Figure 5) that were proposed earlier. The RAIRS experiments with Cl₂ and water at 180 K, which were alluded to previously,¹ gave evidence of the formation and deposition of a reaction product at that temperature. Our experiments, with Cl₂ and H₂O both slowly warming from the condensed phase, indicate that a reaction may begin at temperatures as low as 130 K. In the

stratosphere, the partial pressure of gaseous Cl₂ is typically somewhat lower than the pressures on the order of 10⁻⁷ mbar used in these experiments. However, there have been questions concerning the possible fates of molecular chlorine that may be formed during heterogeneous reactions on solid PSC's. Our experiments indicate that, in addition to desorbing as the molecular species, Cl₂ can react with water to form a condensed film rich in ionic, and potentially reactive, species.

Summary and Conclusions

The interactions of molecular chlorine, dichlorine monoxide, and hypochlorous acid with water ice have been examined under various conditions. At 90–100 K, Cl₂ can be made to condense as a molecular solid on top of a previously condensed ice film. Interaction between the two on the molecular level is minimal. As the temperature is slowly raised, the Cl₂ and water react to produce a mixture of water, hypochlorous acid, and solvated HCl. The reaction begins by about 130 K and, thus, is well underway at the higher temperatures of the stratosphere.

There is evidence that dichlorine monoxide interacts with ice even at 90 K by forming hydrogen bondlike interactions with the H₂O molecules. This interrupts the long-range tetrahedrally coordinated structure of pure hydrogen-bonded water molecules and is manifested by a marked and significant change in the distribution of water-cluster peaks in the positive-ion SIMS spectrum. The change was unique to dichlorine monoxide, a

fact that was attributed to the similarities between the H₂O and Cl₂O molecules. Each has an X₂O structure, with the oxygen atom carrying a δ⁻ charge balanced by a corresponding δ⁺ charge on the X atoms (X = H or Cl). At temperatures of 120 K or higher, the Cl₂O and water react to produce a mixture of hypochlorous acid and water.

References and Notes

- (1) Banham, S. F.; Horn, A. B.; Koch, T. G.; Sodeau, J. R. *Faraday Discuss.* **1995**, *100*, 321.
- (2) Burkholder, J. B. *J. Geophys. Res.* **1993**, *98*, 2963.
- (3) Vickerman, J. C. *Analyst* **1994**, *119*, 513.
- (4) Haynes, D. R.; Tro, N. J.; George, S. M. *J. Phys. Chem.* **1992**, *96*, 8502.
- (5) Sodeau, J. R.; Horn, A. B.; Banham, S. F.; Koch, T. G. *J. Phys. Chem.* **1995**, *99*, 6258.
- (6) Donsig, H. A.; Vickerman, J. C. *J. Chem. Soc., Faraday Trans.* **1997**, *93*, 2755.
- (7) Cady, G. H. *Inorg. Synth.* **1957**, *5*, 156.
- (8) Lancaster, G. M.; Honda, F.; Fukuda, Y.; Rabalais, J. W. *J. Am. Chem. Soc.* **1979**, *101*, 1951.
- (9) Sunner, J.; Ikonomou, M. G.; Kebarle, P. *Int. J. Mass Spectrom. Ion. Processes* **1987**, *82*, 221.
- (10) Donsig, H. A.; Vickerman, J. C. Unpublished data.
- (11) Garvey, J. F.; Herron, W. J.; Vaidyanathan, G. *Chem. Rev.* **1994**, *94*, 1999.
- (12) Echt, O.; Kreisle, D.; Knapp, M.; Recknagel, E. *Chem. Phys. Lett.* **1984**, *108*, 401.
- (13) Stace, A. J.; Moore, C. *Chem. Phys. Lett.* **1983**, *96*, 80.
- (14) Searcy, J. Q.; Fenn, J. B. *J. Chem. Phys.* **1975**, *61*, 5282.
- (15) Donsig, H. A.; Herridge, D.; Vickerman, J. C. To be published.