

Decay Constant for Spontaneous Fission of U^{238} †

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Uranium impurities in minerals undergo spontaneous fission over geologic times, leaving radiation-damage trails whose number can be related to the mineral age, provided λ_F , the spontaneous-fission decay constant of U^{238} , is known. By requiring that the ages of a large number of minerals measured in this way agree with ages determined by decay of K^{40} and Rb^{87} , a value $\lambda_F = 6.9 \times 10^{-17} \text{ yr}^{-1}$ was deduced. By placing a sheet of natural uranium next to a sheet of mica for 6 months and then counting tracks entering the mica from the uranium, a value $\lambda_F = 6.6 \times 10^{-17} \text{ yr}^{-1}$, in good agreement with the above value, was obtained. A weighted average of these two results gives the value $\lambda_F = (6.85 \pm 0.20) \times 10^{-17} \text{ yr}^{-1}$.

THE necessity for making an accurate determination of the decay constant for spontaneous fission of U^{238} arose as a result of the recent discovery that various types of mica,^{1,2} natural glasses,³ and other minerals⁴ contain trails of radiation-damaged material produced by fragments of uranium atoms which have spontaneously fissioned. By chemical etching, these "fossil tracks" can be enlarged to a convenient size for viewing in an optical microscope.^{5,6} In mica the track density ρ_s per cm^2 of mineral surface is related to the time T , since solidification of the rock follows the relation

$$\rho_s = [\exp(\lambda_D T) - 1] \lambda_F N R C_{238} / \lambda_D, \quad (1)$$

where λ_D and λ_F are the total decay constant and spontaneous fission decay constant for U^{238} , N is the number of atoms per cm^3 , C_{238} is the fraction of these atoms that are U^{238} , and R is an effective etchable range of fission fragments in the mineral. Provided T is less than $\sim 10^9 \text{ yr}$, (1) reduces to

$$\rho_s \simeq \lambda_F T N R C_{238}. \quad (2)$$

The necessity for determining N , R , and C_{238} can be avoided by exposing the mineral to a dose n of thermal neutrons and then measuring the density ρ_I of new tracks resulting from thermal fission of U^{235} ,

$$\rho_I = n \sigma N R C_{235}, \quad (3)$$

where σ is the cross section for thermal fission of U^{235} . The mineral age is then simply

$$T = (\rho_s / \rho_I) n \sigma I / \lambda_F, \quad (4)$$

with $I = C_{235} / C_{238}$.

In recent studies of mineral ages by this technique,^{3,7,8} all of the parameters except λ_F were either known or measured to within 10% accuracy. Previously measured values of λ_F ⁹⁻¹³, which appear to be most reliable, range from $\sim 1.2 \times 10^{-16}$ to $\sim 5.3 \times 10^{-17} \text{ yr}^{-1}$. In order to make mineral age determinations on an *absolute* basis by the fission track method, λ_F must also be known to at least 10% accuracy.

In Fig. 1, which uses data, some of which is from our recent paper,³ we have compared our fission track ages of a group of micas⁷ and natural glasses³ (tektites) with age determinations based on the radioactive decay of K^{40} or Rb^{87} , whose decay constants are accurately known.¹⁴ The fission track ages were calculated assuming $\lambda_F = 8.27 \times 10^{-17} \text{ yr}^{-1}$. Taking the straight line drawn through the points as the best fit and requiring that the ages obtained by the two methods be concordant, we deduce a value

$$\lambda_F = (6.9 \pm 0.2) \times 10^{-17} \text{ yr}^{-1}. \quad (5)$$

We now describe an independent experimental determination of λ_F which confirms the above value and permits us to use the age equation (4) on an absolute basis. The method is analogous to the fission-track method of dating minerals.

A sheet of synthetic mica, initially containing no tracks because of its young age, was placed next to a sheet of natural uranium and used to count spontaneous and induced fission fragments emerging from the uranium. The mica-uranium sandwich was cut in two. One-half was stored for six months to collect spontaneous fission events; the other half was irradiated in the

⁷ R. L. Fleischer, P. B. Price, E. M. Symes, and D. S. Miller [to be published in *Science* (1963)].

⁸ M. Maurette, P. Pellas, and R. M. Walker [to be published in *Bull. Soc. Franc. Minéral. Cryst.* (1963)].

⁹ N. A. Perfilov, *Zh. Eksperim. i Teor. Fiz.* **17**, 476 (1947).

¹⁰ E. Segre, *Phys. Rev.* **86**, 21 (1952).

¹¹ P. K. Kuroda and R. R. Edwards, *J. Inorg. Nucl. Chem.* **3**, 345 (1957).

¹² P. L. Parker and P. K. Kuroda, *J. Inorg. Nucl. Chem.* **5**, 153 (1958).

¹³ E. K. Gerling, *Radiokhimiya* **1**, 223 (1959).

¹⁴ Two points at the top of the graph show too low a fission-track age. Such a discrepancy for old minerals has been shown³ to result from track fading due to high temperatures, which are more likely to have occurred in very old samples.

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¹ P. B. Price and R. M. Walker, *Nature* **196**, 732 (1962).

² P. B. Price and R. M. Walker, *J. Geophys. Res.* **68**, 4847 (1963).

³ R. L. Fleischer and P. B. Price (to be published in *J. Geophys. Res.* **69**, Jan. 15, 1964).

⁴ R. L. Fleischer and P. B. Price (unpublished results).

⁵ P. B. Price and R. M. Walker, *Phys. Letters* **3**, 113 (1962).

⁶ R. L. Fleischer and P. B. Price, *J. Appl. Phys.* **34**, 2903 (1963).

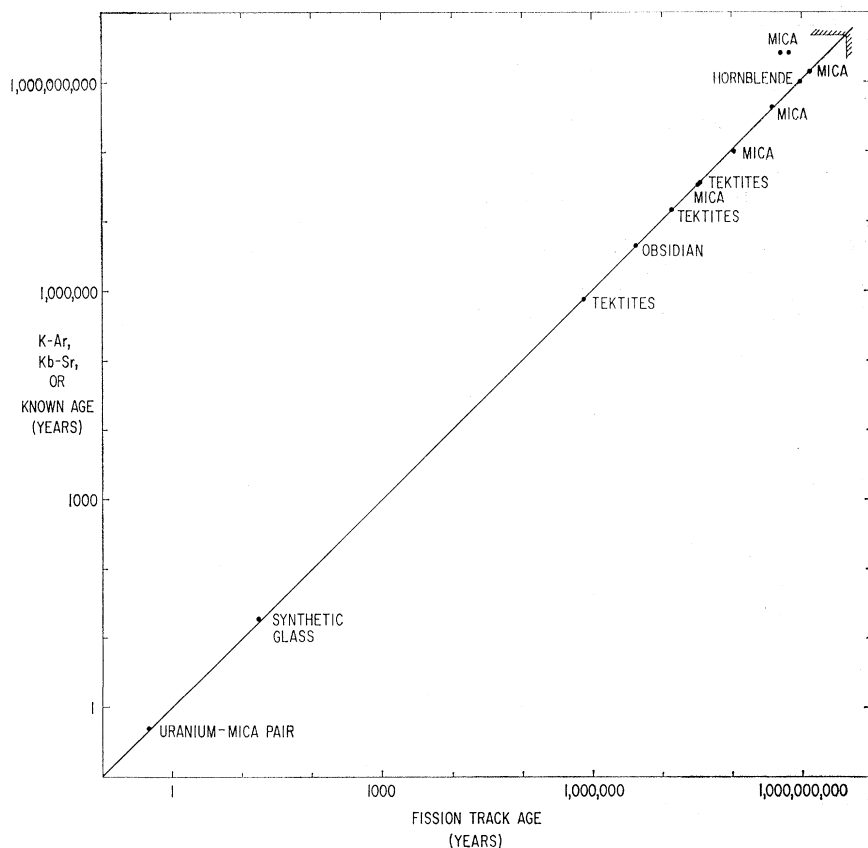


FIG. 1. The graph, which uses data from Ref. 3, unpublished data, and the new experiment reported here, shows that the ages of a large number of micas and glasses, as measured by counting spontaneous-fission tracks, bear a constant relationship to the ages determined by decay of K^{40} and Rb^{87} in the minerals. By drawing a straight line through the points on the log-log plot and equating the ages, a value $\lambda_F = 6.9 \times 10^{-17} \text{ yr}^{-1}$ for the spontaneous-fission decay constant was deduced.

thermal column of the Brookhaven reactor to a dose of 1.04×10^{11} neutrons/cm² to collect induced fission events.

Equation (4) was used to calculate λ_F . The measured values of ρ_s and ρ_I were $(210 \pm 21)/\text{cm}^2$ and $(2.73 \pm 0.19) \times 10^6/\text{cm}^2$, respectively. The exposure time for the spontaneous-fission tracks was 0.51 yr and the best values of σ and I were taken to be $\sigma = (5.82 \pm 0.04) \times 10^{-22} \text{ cm}^2$ and $1/I = 137.8$.^{15,16} Substitution into (4) gave

$$\lambda_F = (6.6 \pm 0.8) \times 10^{-17} \text{ yr}^{-1}, \quad (6)$$

where the indicated standard deviation pertains to counting statistics. This value thus agrees with (5) within the counting errors.

Of the two values (5) and (6), the former, $\lambda_F = 6.9$

¹⁵ E. K. Hyde, University of California, Lawrence Radiation Laboratory Report No. UCRL-9036-Rev., 1962, p. 67 (unpublished).

¹⁶ E. K. Hyde, University of California, Lawrence Radiation Laboratory Report No. UCRL-10612, 1963, p. 37 (unpublished).

$\times 10^{-17} \text{ yr}^{-1}$, is probably more reliable since it is based on a comparison of a large number of experimental measurements of mineral ages ranging from 6×10^5 up to 1.4×10^9 yr (Fig. 1). The only significant sources of error are the decay constants for K^{40} and Rb^{87} , which are known to within 5%,¹⁷ and the thermal neutron dose, which was measured to within $\sim 5\%$ by counting the Ba^{140} and Mo^{99} activity in a U foil. A weighted average of the two results gives $\lambda_F = (6.85 \pm 0.20) \times 10^{-17} \text{ yr}^{-1}$.

The main advantages of the fission-track method over previous methods involving electronic counting, emulsion counting, or radiochemical analysis are its simplicity and its ability to discriminate completely against a background of light particles such as cosmic rays and alpha particles.¹⁸

¹⁷ L. T. Aldrich and G. W. Wetherill, *Ann. Rev. Nucl. Sci.* **8**, 257 (1958).

¹⁸ R. L. Fleischer, E. L. Hubbard, P. B. Price, and R. M. Walker, *Phys. Rev.* (to be published).