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INFLUENCE OF FREEZING - UNFREEZING CYCLES ON THE THERMAL BEHAVIOUR OF A MARINE CLAY

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## **ABSTRACT**

After ten freezing - unfreezing cyclee a marine cley indicated an important change of water sorption /particle thickness/ and of weight loss in static heating up to 800 °C. This is due to changes in aicrostructure during freezing.

# **INTRODUCTION**

Variability in water sorption, WS, is attributed here to change in specific surface, S, and in particle thickness,  $\delta$ . It waa observed in aeveral processes, e.g. mechanical treatment, drying, storage at a high water content etc., see  $[7]$ . These processes may influence as well weight loss in etatic heating.  $A G.$  All these influences depend not only on pretreatment but also on form and size of sample under study. In powdered sample these effects mey disappear, though some change in behaviour of Berkbent bentonite powder, depending on sample pretreatment, wae observed. Also ite microetructure wes different whether it was atored at decreasing or at increasing relative humidity  $[8]$ .

Here a study will be presented of the influence of cyclic freezing on WS and on TG behaviour of a marine clay  $\lceil 6 \rceil$ . A similar influence, i.e. a decrease in WS and increase in particle thickness was observed in a silty leam by Skarżyńska  $\lceil 3 \rceil$ . SEM atudy of this soil indicated some microstructural changes  $\lceil 10 \rceil$ which may occur also in this marine clay. Some phase transfor-Wations are possible as well, during freezing and/or drying and/or heating which are discussed in the adjoining paper  $\lceil 9 \rceil$ .

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#### MATERIALS AND METHODS

Three samples of a marine clay were taken from the Baltic sea bottom at depths of /1/ 3.6 m, /2/ 5.4 m, and /3/ 5.5 m.

Mineral composition of clays from this region was studied previously. They contain montmorillonite, partly interstratified with illite, and keolinite in variable proportions, as well as quartz and feldepar in coarser fractions. Clay fraction may give XRD patterns of an almost amorphous matter, showing better crystallinity after homoionization to Mg or K /Fig. 1/.







- Fig. 1. X-ray diffraction (XRD) of two oriented samples of a marine clay form Baltic sea bottom:
	- a) clay fraction,
	- b) homoionized to K.
		- c) homolonized to Ma.

Geotechnical properties are indicated in Table 1. The natural water content was high,  $W_{n} = 42 - 66$  %. Content of crystal phase water was estimated at  $W_h = 13 - 15$  %, diffuse layer water at  $W_{\text{el}} = 16 - 19$  % and macropore /interaggregate/ water  $W_{\text{max}}$ = 13 - 32 %. Plestic limit was  $W_{\text{p}}$  = 21 - 28 % and liquid limit W<sub>1</sub> = 51 - 69 %. Undrained shear strength was low,  $8_H/U$  -5 - 11 kPa and it decreased after remolding to  $S_{11}/R/ = 2 - 6$  kPa. Thus sensitivity  $S_{11}/U/1S_{11}/R/$  varied between 1.8 and 2.5.

Remolded clays /R/ were placed in Wasiliew apparatus, were wrapped in plestic foil and freezed ten timas for 6-7 h at - 20 <sup>O</sup>C leaving for ca. 18 h to unfreeze to room temperature. Changes in sample height were recorded: after shrinking in the first to third freezing cycle, some awelling was observed which was dependent on sampling depth, i.e. decreased from

TABLE 1 Geotechnical properties of the marine clay

			Nr Depth W W W S kPa Sensi- W W W Ho m $X^0$ $X^0$ $X^1$ $X^1$ $X^2$ $X^2$ $X^3$ $X^4$ $X^5$ $X^6$ $X^7$ $X^8$			
			1 3.6 66 28 69 5 2 2.5 15 19 32 43			
			2 5.4 50 22 51 9 3 3.0 13 16 21 36			
			$3$ 5.5 42 21 52 11 6 1.8		13 16 13 37	

 $/1/$  21  $%$  to  $/2/$  13  $%$  and  $/3/$  6  $%$ . Montmorillonite content  $/$ Mo/ decreased in a similar aequence from 43 % to 37 %.

Grain size distribution was measured before and after cyclic freezing and mainly some increase in content of coarser fractions was observed /Table 2/, especially in sample /3/ taken from the highest depth. This indicates eggregation processes.

TABLE 2

Final increase in volume 4h/h, and change in grain size distribution after freazing



To measure WS three unground specimen of ca. 300 mg of each sample were placed in aluminium containsre, were dried at 110°C end wers stored for 10 days in a desiccator over the saturated Mg/NO<sub>3</sub>/<sub>2</sub> solution /p/p<sub>o</sub>= 0.5/, efter which they were transferred to  $p/p_0$  = 0.95 /over the 10 % H<sub>a</sub>SO<sub>4</sub> solution/ for 16 days. They were dried again at 110 <sup>o</sup>C and heated successively at 220 <sup>o</sup>C, 400 °C and 800 °C. Weight changes were recorded and  $\triangle G/G_0$  was calsulated assuming as G the initial weight of the oven dried specimen /Table 3/.

TABLE 3

Water sorption and weight loss in static heating in percent /mean of three samples/

Nr	Sample Depth	State	$M$ at $p/p_0 =$ = 0.5 %	W at $p/p_0 =$ = 0.95 %	$\triangle G$ /100- - 800 °C	W/0,95/ $- \triangle G$ $\boldsymbol{\mathsf{x}}$
1	$3.6 -$ $-3.7$ $\blacksquare$	U R F	2,9 $+ 0.2$ $2.5 -$ $^{\circ}$ 0.2 1.6 0.3	15.2 0.5 $\ddot{\phantom{1}}$ $16.6 - 1.3$ 10.0 1.6	$-10.8$ 0.2 $\ddot{\phantom{1}}$ $-11.6$ 0.7 $-15.4$ 0.8	26 28 25
2.	$5.4 -$ $-5.5$ $\blacksquare$	U R F	0.1 1.7 ٠ 1.8 $-0.2$ 1.4 0.2	$12.8 + 9.8$ 11.8 $-1.0$ 10.5 0.3	$-14.3 +$ 0.4 $-14.5 - 0.5$ $-15.1$ 1.5	27 26 26
3.	$5.5 -$ $-5.6$	υ R F	1.8 0.2 $\bullet$ 2.0 0.3 1.9 0.2	13.0 1.5 $\bullet$ 13.2 0.5 1.9 11.3	$-15,0$ $+0.04$ $-15.0$ 0.3 $-11.6$ 0.9	28 28 23

 $U -$  undisturbed,  $R -$  remolded,  $F -$  frozen cyclically

**DT SCUSS TON** 

All the data /XRD, WS, TG/ indicate that marine clavs from the given region have a very unstable particle thickness, especially at lower sampling depth /sample 1/. WS at high  $p/p_{n} = 0.95$ is much bigger than that at  $p/p_0 = 0.5$ . Their ratio has an exceptionally important value of 5.2 to 7.5, whereas in case of a constant particle thickness it is ca. 1.5.

Prolonged storage in dry state caused a decrease in WS /thus particle collapse/ by ca. 50  $\frac{1}{2}$  /  $\frac{1}{2}$  . Here freezing has a similar effect, also dependent on sampling depth. Relative change in WS varied from 36-40 % /1/ to 26-11 % /2/ end 4-14 % /3/ for  $p/p_{n}$ = = 0.5 and 0.95 reapectively.

It was found that around ice lenses there is formed a parallel particle arrangement both in monomineral clays  $\begin{bmatrix} 2 \end{bmatrix}$  and in soils  $\lceil$ 10]. This increase in domain thickness is due to /1/ cryosuction  $\begin{bmatrix} 1 \\ 1 \end{bmatrix}$  and /2/ pressure exerted by the ice crystal being formed. Such a microstructure may have an influence on WS, on temperature of dehydration and on that of dehydroxylation as well as on phase transformation during heating, sse [9]. If unrestricted, the perallel atructure may sorb water and awall fre- $\ddot{\mathbf{a}}$ ly, thus high particle delamination may occur  $\begin{bmatrix} 4 \end{bmatrix}$ . If trapped within the cluster structure /Fig. 2/ its water scrption and ewelling is restricted, water escape during heating and dehy-

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droxylation is hindered and phase transformation is most probable. Indeed decreased water acrption and important differences in  $\Delta$  G were observed during heating /Table 3/.

After WS measurement the weight of the oven dried specimen decreased by 0.1 to 0.4 % =  $\triangle 6/110$  °C/. After heating at 220 °C it decreased only by 0.8 to 1  $\chi$ , whereas sample pretreatment influenced  $\triangle$  G/400  $^{\circ}$ C/, which varied between 1.5 and 2.7  $\chi_4$  even more the  $\triangle$  G/800  $^{\circ}$ C/ was influenced /11 to 15 %/, which change was just opposite to that of W /O.95/. To show this tendency the sum of W/O.95/ and -  $\triangle$  G/800  $^{\circ}$ C/ is presented in the last column of Table 3. This total change in weight varies in undisturbed samples and in the remolded ones between 26 and 28 %, whereas in freezed samples this variability ranges between 23 and 25.6 %. Thus the smallest total change in weight was indicated by freezed samples, where much parallel particle agrangement is to be expected. In this case not only increase in temperature of dehydration may occur, but also phase transformation is possible. Berkbent bentonite of parallel atructure did not dehydrate completely at 400  $^{\circ}$ C /and posaibly even at 800  $^{\circ}$ C/, which was indicated by a broadened peak at 9.7  $\frac{2}{3}$ . This was possible due to surface melting of the parallel structure at 800  $^{\circ}$ C.

Heating of bentonite resulted in phase transformation, which was indicated by new XRD peaks, that were mest abundant in parallel structure, with cristoballite prevailing, whereas in cluster structure feldspar could be found. Depending on these transformations weight loss in heating may be either increased or decreasad.

It may be concluded that suitable study of water sorption and weight loss in static heating of the pretreated apecimen may give eome information on changes in microstructure during the given process. Grinding or pulverization of the apecimen may annihilate the investigated effect.

# **SUMMARY**

In marine clays water eorption and weight loss in static heating change after cyclic freezing. The most probable explanation is the development in freezing of parallel structure /face--to-face contacts/. At lower sampling depth this causes mostly a decrease in WS and increase in temperature of dehydration. At higher sampling depth the additional effect may be the phase



Fig. 2. Parallel atructure formed within cluster matrix of bentenite suspensions:/a/ Brebent, B/R-2, 500x, picture<br>breedth 200 um, /b/ Berkbent, B/B-A, 2000x, picture<br>breedth 50 um, 9 breadth 50 um

transformation, which causes a decrease in  $\triangle$  G /800  $^{\circ}$ C/.

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