INFLUENCE OF FREEZING - UNFREEZING CYCLES ON THE THERMAL BEHAVIOUR OF A MARINE CLAY

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#### **ABSTRACT**

After ten freezing - unfreezing cycles a marine cley indicated an important change of water sorption /particle thickness/ and of weight loss in static heating up to 800 °C. This is due to changes in microstructure during freezing.

## INTRODUCTION

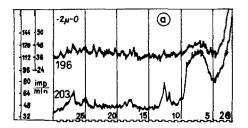
Variability in water sorption. WS, is attributed here to change in specific surface, S, and in particle thickness,  $\delta$ . It was observed in several processes, e.g. mechanical treatment, drying, storage at a high water content etc., see [7]. These processes may influence as well weight loss in static heating,  $\Delta$  G. All these influences depend not only on pretreatment but also on form and size of sample under study. In powdered sample these effects may disappear, though some change in behaviour of Berkbent bentonite powder, depending on sample pretreatment, was observed. Also its microetructure was different whether it was stored at decreasing or at increasing relative humidity  $\lceil 8 \rceil$ .

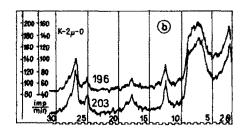
Here a study will be presented of the influence of cyclic freezing on WS and on TG behaviour of a marine clay  $\begin{bmatrix} 6 \end{bmatrix}$ . A similar influence, i.e. a decrease in WS and increase in particle thickness was observed in a silty leam by Skarżyńska  $\begin{bmatrix} 3 \end{bmatrix}$ . SEM atudy of this soil indicated some microatructural changes  $\begin{bmatrix} 10 \end{bmatrix}$  which may occur also in this marine clay. Some phase transformations are possible as well, during freezing and/or drying and/or heating which are discussed in the adjoining paper  $\begin{bmatrix} 9 \end{bmatrix}$ .

## MATERIALS AND METHODS

Three samples of a marine clay were taken from the Saltic see bottom at depths of /1/3.6 m, /2/5.4 m, and /3/5.5 m.

Mineral composition of clays from this region was studied previously. They contain montmorillonite, partly interstratified with illite, and keelinite in variable proportions, as well as quartz and feldapar in coarser fractions. Clay fraction may give XRD patterns of an almost amorphous matter, showing better crystallinity after homoionization to Mg or K /Fig. 1/.





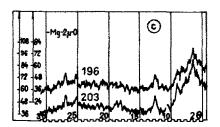


Fig. 1, X-ray diffraction (XRD) of two oriented samples of a marine clay form Baltic sea bottom:

- a) clay fraction,
- b) homoionized to K,
  - c) homoionized to Mg.

Geotechnical properties are indicated in Table 1. The natural water content was high,  $W_{\rm n}=42-66$  %. Content of crystal phase water was estimated at  $W_{\rm h}=13-15$  %, diffuse layer water at  $W_{\rm d}=16-19$  % and macropore /interaggregate/ water  $W_{\rm mac}=13-32$  %. Pleatic limit was  $W_{\rm p}=21-28$  % and liquid limit  $W_{\rm l}=51-69$  %. Undrained shear strength was low,  $8_{\rm l}/{\rm l}/\approx5-11$  kPa and it decreased after remolding to  $S_{\rm l}/{\rm R}/\approx2-6$  kPa. Thus sensitivity  $S_{\rm l}/{\rm l}/{\rm l}/{\rm s}_{\rm l}/{\rm R}/$  varied between 1.8 and 2.5.

Remolded clays /R/ were placed in Wasiliew apparatus, were wrapped in plastic foil and freezed ten times for 6-7 h at - 20 °C leaving for cs. 18 h to unfreeze to room temperature. Changes in sample height were recorded: after shrinking in the first to third freezing cycle, some swelling was observed which was dependent on sampling depth, i.e. decreased from

TABLE 1					
Geotechnicel	properties	of	the	merine	cley

Nr	Depth	₩ % <sup>n</sup>	₩ <sub>P</sub>	X,	S	kPe R	Sensi- tivity	%h	%d	Xaac	Mo %
1	3.6	66	28	69	5	2	2,5	15	19	32	43
2	5.4	50	22	51	9	3	3.0	13	16	21	36
3	5.5	42	21	52	11	6	1.8	13	16	13	37

/1/ 21 % to /2/ 13 % and /3/ 6 %. Montmorillonite content /Mo/ decreased in a similar sequence from 43 % to 37 %.

Grain size distribution was measured before and after cyclic freezing and mainly some increase in content of coarser fractions was observed /Table 2/, especially in sample /3/ taken from the highest depth. This indicates eggregation processes.

TABLE 2
Final increase in volume 4h/h, and change in grain size distribution after freezing

Nr	Depth m	4 h/h %	Content of fraction %			
			eend	eilt	cley	
			0,05-2 mm	0,002-0,05	ma < 2/48	
1	3,6±3,7 after freezing	21	6 1	3€ 41	58 58	
2	5,4+5,5 after freezing	13	19 20	31 26	50 54	
3	5,545,6 after freezing	6	12 26	34 24	54 50	

To measure WS three unground specimen of ca. 300 mg of each sample were placed in aluminium containers, were dried at  $110^{\circ}$ C and were stored for 10 days in a desiccator over the saturated Mg/NO<sub>3</sub>/<sub>2</sub> solution /p/p<sub>0</sub>= 0.5/, after which they were transferred to p/p<sub>0</sub> = 0.95 /ower the 10 % H<sub>s</sub>SO<sub>4</sub> solution/ for 16 days. They were dried again at 110 °C and heated successively at 220 °C, 400 °C and 800 °C. Weight changes were recorded and  $\Delta$  G/G<sub>0</sub> was calculated assuming as G<sub>0</sub> the initial weight of the oven dried specimen /Table 3/.

TABLE 3
Water sorption and weight loss in static heating in percent
/mean of three samples/

Nr	Sample Depth	State	W at p/p = - 0.5 %	W at p/p= = 0.95 %	△ G /100- - 800 °C %	W/0,95/ -△G %
1	3.6 - - 3.7 m	U R F	$\begin{array}{c} 2.9 \pm 0.2 \\ 2.5 & 0.2 \\ 1.6 & 0.3 \end{array}$	15.2 ± 0.5 16.6 1.3 10.0 1.6	-10.8 ± 0.2 -11.6 0.7 -15.4 0.8	26 28 25
2	5.4 - - 5.5	U R F	1.7 ± 0.1 1.8 0.2 1.4 0.2	12.8 ± <b>6</b> .8 11.8 1.0 10.5 0.3	-14.3 ± 0.4 -14.5 0.5 -15.1 1.5	27 26 26
3	5.5 - - 5.6	U R F	1.8 ± 0.2 2.0 0.3 1.9 0.2	13.0 ± 1.5 13.2 0.5 11.3 1.9	-15.0 ± 0.04 -15.0 0.3 -11.6 0.9	28 28 23

U - undisturbed, R - remolded, F - frozen cyclically

#### DISCUSSION

All the data /XRD, WS, TG/ indicate that marine clays from the given region have a very unstable particle thickness, especially at lower sampling depth /sample 1/. WS at high  $p/p_0 = 0.95$  is much bigger than that at  $p/p_0 = 0.5$ . Their ratio has an exceptionally important value of 5.2 to 7.5, whereas in case of a constant particle thickness it is ca. 1.5.

Prolonged storage in dry state caused a decrease in WS /thus particle collapse/ by ca. 50 %  $\begin{bmatrix} 7 \end{bmatrix}$ . Here freezing has a similar effect, also dependent on sampling depth. Relative change in WS varied from 36-40 % /1/ to 26-11 % /2/ and 4-14 % /3/ for p/p<sub>0</sub>= = 0.5 and 0.95 respectively.

It was found that around ice lenses there is formed a parallel particle arrangement both in monomineral clays [2] and in soils [10]. This increase in domain thickness is due to /1/ cryosuction [1] and /2/ pressure exerted by the ice crystal being formed. Such a microstructure may have an influence on WS, on temperature of dehydration and on that of dehydroxylation as well as on phase transformation during heating, see [9]. If unrestricted, the parallel atructure may sorb water and swell fresly, thus high particle delamination may occur [4]. If trapped within the cluster structure /Fig. 2/ its water sorption and swelling is restricted, water escape during heating and dehy-

droxylation is hindered and phase transformation is most probable. Indeed decreased water sorption and important differences in  $\Delta$  G were observed during heating /Table 3/.

After WS measurement the weight of the oven dried apecimen decreased by 0.1 to 0.4 % =  $\Delta G/110^{\circ}$ C/. After heating at 220 °C it decreased only by 0.8 to 1 %, whereas sample pretreatment influenced  $\triangle$  G/400 °C/. which varied between 1.5 and 2.7 %; even more the A G/800 °C/ was influenced /11 to 15 %/, which change was just opposite to that of W /0.95/. To show this tendency the sum of W/O.95/ and -  $\Delta$  G/800  $^{
m O}$ C/ is presented in the last column of Table 3. This total change in weight varies in undisturbed samples and in the remolded ones between 26 and 28 %, whereas in freezed samples this variability ranges between 23 and 25.6 %. Thus the smallest total change in weight was indicated by freezed samples, where much parallel particle agrangement is to be expected. In this case not only increase in temperature of dehydration may occur, but also phase transformation is possible. Berkbent bentonite of parallel structure did not dehydrate completely at 400 °C /and possibly even at 800 °C/, which was indicated by a broadened peak at 9.7 % . This was possible due to surface melting of the parallel structure at 800 °C.

Heating of bentonite resulted in phase transformation, which was indicated by new XRD peaks, that were most abundant in parallel atructure, with cristoballite prevailing, whereas in cluster atructure feldsper could be found. Depending on these transformations weight loss in heating may be either increased or decreased.

It may be concluded that suitable study of water corption and weight loss in static heating of the pretreated apecimen may give come information on changes in microstructure during the given process. Grinding or pulverization of the specimen may annihilate the investigated effect.

#### SUMMARY

In marine clays water corption and weight loss in static heating change after cyclic freezing. The most probable explanation is the development in freezing of parallel structure /fece-to-face contacte/. At lower sampling depth this causes mostly a decrease in WS and increase in temperature of dehydration. At higher sampling depth the additional effect may be the phase

a



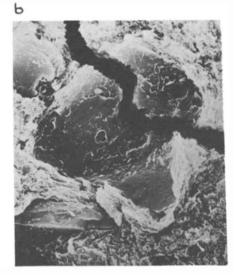


Fig. 2. Parallel atructure formed within eluster matrix of bentenite suspensions:/a/ Brebent, B/R-2, 500x, picture breadth 200 µm, /b/ Berkbant, B/B-A, 2000x, picture breadth 50 µm 9

transformation, which causes a decreese in A G /800 °C/.

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